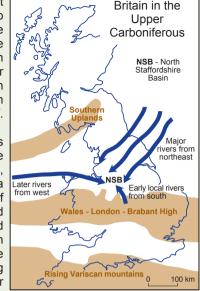
GEOLOGY GEOLOGY LOCATION MAP

The rocks of the Churnet Valley belong to two geological time periods, the Upper Carboniferous (here about 318 - 313 million vears old) and the Lower Triassic (250 - 245 million years old). See the column next to the map overleaf.

Carboniferous sediments were deposited near to the equator, close to sea-level. in a low-lying muddy delta on the southern margin of a continent called Laurussia. Muds. silts and sands were laid down in cycles due to global climate change. Global warming melted the southern polar



ice cap. Sea-level then rose, flooding the delta depositing muds containing marine fossils such as goniatites. Global cooling then caused the ice caps to grow again, sea-level

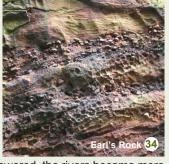
Froghall Ironstone

fell and rivers deposited muds and then sandier material in the delta. Eventually the delta surface emerged. Plants colonising this new land formed peat in swamps. Later this was buried and turned into coal. The lower part of the sequence contains thicker and coarser sand units forming part of the Namurian 'Millstone Grit' sequence; the upper muddier part

containing more coal and ironstone belongs to the Westphalian 'Coal Measures'.

Late in the Carboniferous Laurussia collided with the continent of Gondwana to the south to form the supercontinent Pangaea. The rocks were then uplifted and folded in a mountain building episode called the Variscan

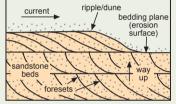
Rocks of Permian age are missing here due to that uplift. During the Triassic, the region was at northern hot desert latitudes. Large braided rivers deposited sands and pebbles unconformably over the eroded Carboniferous rocks (see unconformity diagram below). These rivers flowed from the fasteroding Variscan mountains in south-west England and Brittany.



As the highlands were gradually lowered, the rivers became more meandering and the sands finer grained. The rocks were later fractured and faulted by late Triassic earth movements associated with the opening of the Atlantic Ocean.

### **CROSS BEDDING**

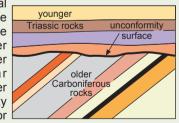
Cross-bedding is a feature found within lavered (bedded) sediments and sedimentary rocks. It consists of a series of sloping layers between the main. normally flat, bedding plane surfaces. It is formed by the deposition of sediment on the



downcurrent side of dunes and ripples in either air or water. Erosion of the tops of these sloping layers can be used to determine the 'way up' of sedimentary rocks, and the slope direction can be used to determine the direction of the flow of ancient rivers and winds.

### UNCONFORMITY

An unconformity is a geological surface which represents a time gap in the geological record. The surface separates a younger sequence above from an older one below. An angular unconformity is where the older rocks below the unconformity have been uplifted, tilted and/or



folded and then exposed to erosion before the deposition of the younger ones.

In the Churnet Valley the Carboniferous rocks were uplifted and tilted gently westwards at the end of the Carboniferous Period during the Variscan Orogeny. The sequence was then eroded down to the Lower Westphalian / Upper Namurian level during the Permian and deposition resumed in the Lower Triassic above the unconformity. Thus, this unconformity represents about a 65 million year time gap.

## Bakewell Leek. Buxton & Macclesfield T •••• GEOTRAIL Stokeon-Trent Ashbourne upon Trent

A project managed by Richard Waller and lan Stimpson on behalf of the Staffordshire RIGS Group with financial support from the Staffordshire Aggregates Levy Grant Scheme 2007. Text and photos by Patrick Cossey, Peter Floyd, John Reynolds, Ian Stimpson and Richard Waller unless otherwise stated. Design by Rosie Duncan.

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A member of UKRIGS, the Staffordshire RIGS Group is a voluntary organisation responsible for the conservation of regionally important geological and geomorphological sites in Staffordshire. For more information contact: ffordshire RIGS - www.staffs-rigs.org.uk

North Staffordshire Group Geologists' Association - www.esci.keele.ac.uk/nsgga Staffordshire Wildlife Trust - www.staffs-wildlife.org.uk

This trail is dedicated to the memory of Ken Rout, founding member of SRIGS.

Remember to follow the country code and please do not hammer the rock surfaces. Be safe, plan ahead and follow any signs; leave gates and property as you find them; protect plants and animals and take your litter home. Keep dogs under close control and consider other people.

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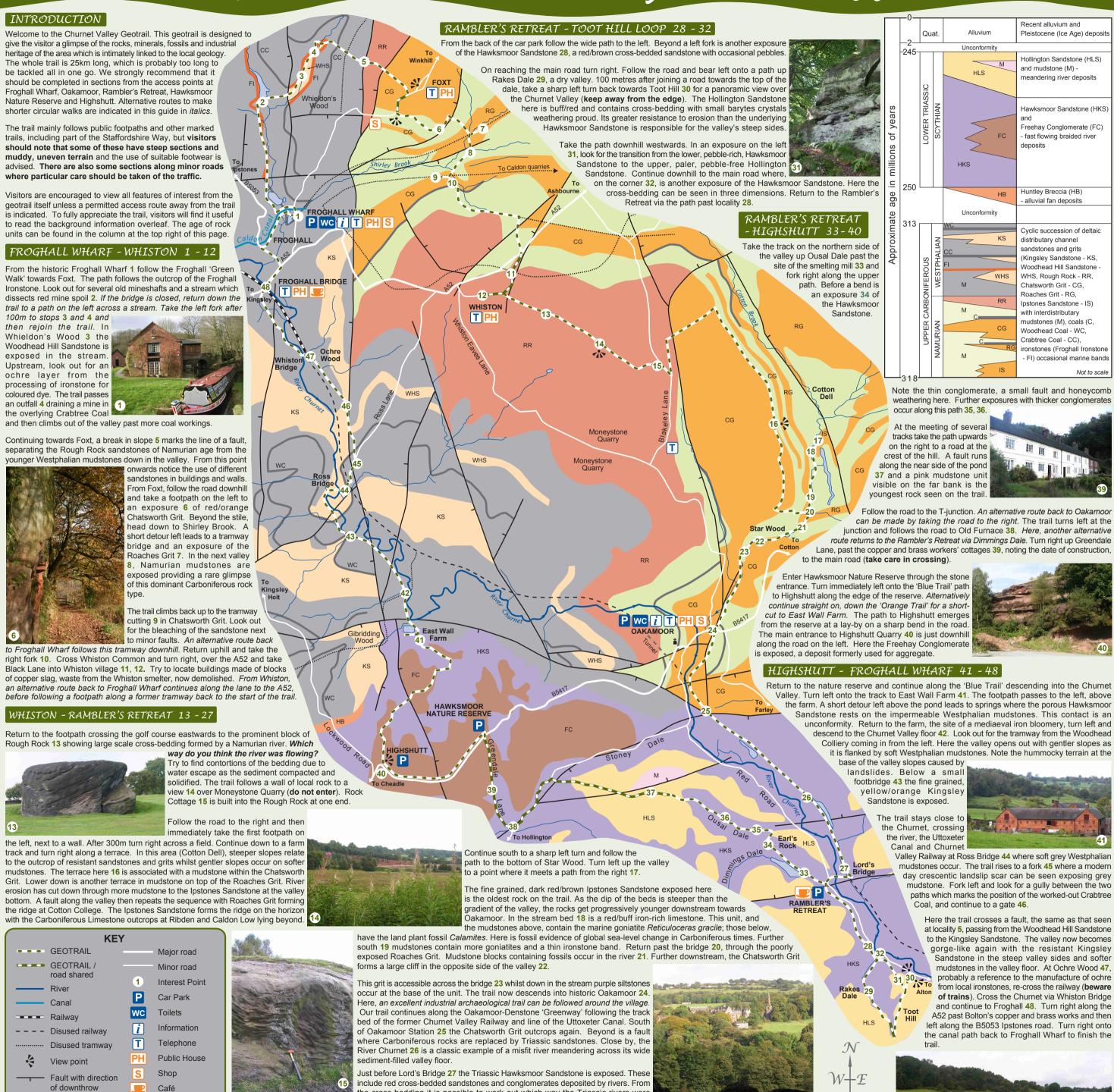


# The Churnet Valley Geotrail

**True Grits and Ghosts of** a Great Industrial Past



# The Churnet Valley Geotrail



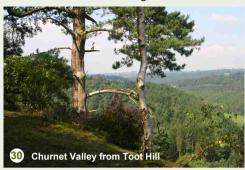
the cross-bedding it is possible to work out which way the Triassic rivers were

flowing. How does this compare with the Carboniferous river deposits at 13 and the modern Churnet? Cross

Lord's Bridge towards the Rambler's Retreat.

The Churnet Valley is renowned for its steep-sided and dramatic character and this has led to the area being nicknamed "Little

Switzerland". Today, its sheer size is at odds with the placid misfit river meandering across its picturesque floor. However, during the last Ice Age 18,000 years ago, part of the British



ice sheet penetrated the upper part of the Churnet Valley. Torrents of meltwater flowing under the ice and issuing from the glacier margin caused huge amounts of erosion and sediment transport, creating the valley's present form. Changes in the valley's width and depth reflect variations in the resistance of the underlying



rocks to erosion. The gorge-like sections pass through resistant sandstones and more open areas are floored by softer mudstones. The ice sheet also caused a complete reversal of the river's direction. Previously, the Churnet flowed north from Froghall to

the north of Cheddleton before turning west and joining the Trent drainage system at Endon. When ice blocked this route it forced the river to flow southward and erode its present course.

Incision and the creation of the valley's steep sides led to numerous landslides, particularly in areas underlain by weak mudstones. Down-cutting of the main valley also "rejuvenated" the tributary valleys as they attempted to erode down to the new lower level of the Churnet. Their erosion was not caused by glacial meltwater,

but by the action of seasonal snowmelt streams. These flowed across the frozen ground in front of the ice sheet and left a series of dry valleys, which currently have no surface streams.



19th centuries the Churnet Valley was the scene of intense industrial activity. with furnaces. forges, mills and mines. The underlying bedrock of the valley contained



ores of iron, copper and lead, as well as coal, limestone and sandstone. The River Churnet provided water power and the woods provided charcoal. Tramways were constructed to carry limestone from the Caldon Low quarries and local coal and ironstone to the Caldon (1778) 1 and Uttoxeter (1809) 1, 24-27, 44 canals. The Churnet Valley Railway (1849) 24-27, 44, 47, 48 transported industrial sand and copper products from Oakamoor onto the rail network.

INDUSTRIAL HERITAGE



A medieval iron bloomery existed at East Wall 41 in the 12th-14th centuries. Ironstone from the local Coal Measures was mixed with charcoal and limestone, covered with turf and heated to separate the iron. On cooling, the slag was raked from the iron

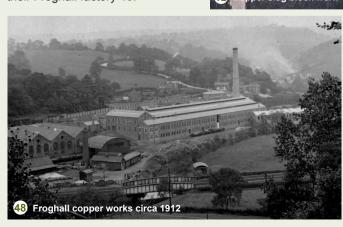
mass (or pig). An iron forge (hammer mill), using water power, is recorded at Oakamoor 24 in 1573. One of the first iron furnaces equipped with bellows geared from a water wheel was built in 1593 at Old Furnace 38. The Froghall Ironstone seam was mined around Froghall where spoil is still visible 2.





smelting. Cheadle Brass & Copper Co., was set up in 1734. Extensive mining of copper ore at Ecton in the Manifold Valley required building a new copper smelting works at Whiston (1768) 11. Blocks of copper slag can be

seen in local buildings and walls 11, 12. By 1828 most copper wire production was handled at the new mills at Oakamoor 24. A row of terraced cottages built for brass workers can be seen at Greendale 39. By 1834 all copper wire interests had been bought by Thomas Bolton & Son Ltd. of Birmingham and produced at their Froghall factory 48.



For a short time in the mid 1700s, lead mined at Ribden, near Farley, was smelted at the smelting mill in Ousal Dale 33.



In 1775 the Caldon Low quarries were leased to Caldon Lime Co. A tramway from the quarries to the Caldon Canal 1 at Froghall Wharf was part of the Canal Act. Several tramway lines were built (1778) **7**, (1785) **10**, (1804) 11 and (1847) 9. In many places lime kilns can be seen as at Froghall Wharf 1 and Oakamoor 24. These were used to burn

limestone to make lime for use in agriculture and construction. The limestone was and continues to be used as a flux in iron smelting and, with mudstone, in the manufacture of cement.

## **Sandstone and Conglomerates**

Sand has been extracted from the Rough Rock at Moneystone 14 for glassmaking and the Freehay

Conglomerate quarried for aggregates at Highshutt 40 and Croxden. The sandstones are also used widely in the building industry. The

Hollington Sandstone for example is used in many Staffordshire buildings including Pugin's St. Giles Church at Cheadle (1846).



